



Preface

Diffusion and partitioning in planetary interiors

3 1. Introduction

4 This special issue gives an overview of recent
5 progress in the related fields of element partitioning
6 and diffusion at high pressures and temperatures. Geo-
7 chemical models of large-scale differentiation pro-
8 cesses in planetary interiors, such as core formation
9 and partial melting in magma oceans or beneath mid-
10 ocean ridges, require accurate data on how major
11 and trace elements distribute themselves between the
12 relevant phases (including solid and molten silicates,
13 oxides and metal alloys) at high pressures and tem-
14 peratures. Major and trace element diffusion data are
15 important for evaluating whether chemical equilib-
16 rium is likely to be achieved during these differentia-
17 tion processes, and to quantify the extent of chemical
18 interaction that can be expected over time. Diffusion
19 rates are also fundamentally linked to several im-
20 portant physical properties of minerals and liquids,
21 including rheology, and thus contribute to our under-
22 standing of dynamical processes such as convection
23 in planetary interiors.

24 Two conferences held in the 1970s (the Confer-
25 ence on Geochemical Transport and Kinetics, held
26 in June 1973, and the International Conference on
27 Experimental Trace Element Geochemistry ('Sedona
28 Conference'), held in September 1977), can be seen as
29 the starting point for geochemical research on diffu-
30 sion and trace element partitioning in phases relevant
31 to planetary interiors. Since then, novel experimental
32 and analytical techniques have allowed experimental
33 measurements of diffusion and partitioning at increas-
34 ing elevated pressure and temperature conditions, of
35 increasingly direct relevance to processes taking place
36 in the interiors of planets.

37 This special issue brings together 11 contributions
38 spanning the pressure range from 1.5 to 24 GPa, the

latter equivalent to the top of the lower mantle on 39
Earth, and close to the Martian core–mantle bound- 40
ary. We hope that the bringing together of papers on 41
diffusion and partitioning will lead to increased col- 42
laboration between researchers in these two fields and 43
lead to new, more stringent constraints on large-scale 44
geochemical differentiation processes at elevated pres- 45
sures and temperatures. 46

2. Brief summary of contributions 47

Béjina et al. review models for the variation of 48
diffusion rates with pressure, and discuss currently 49
available data sets on diffusion in minerals at elevated 50
pressures. They show the immense progress that has 51
been made in experimental and analytical techniques, 52
particularly during the past 5 years, but also highlight 53
discrepancies that exist between different measure- 54
ments of diffusion at high pressure. As a further il- 55
lustration of recent advances, Holzapfel et al. present 56
new measurements of the pressure and temperature 57
dependence of Fe–Mg inter-diffusion in $(\text{Mg}_x\text{Fe}_{1-x})\text{O}$ 58
ferropericlast, an important phase in Earth's lower 59
mantle, as a function of iron content. They show 60
an emerging consensus on the effect of pressure on 61
cation diffusion rates in this phase, and use their data 62
to constrain the extent of chemical interaction be- 63
tween Earth's core and lower mantle since the time of 64
core formation. Koga et al. use recently obtained data 65
for carbon diffusion in diamond, in conjunction with 66
nitrogen diffusion data, to model the decay of isotopic 67
heterogeneities in diamond as a function of residence 68
time and temperature in the Earth's mantle. The pa- 69
per illustrates the application of diffusion data to the 70
subject of chronometry, and is an example of the type 71
of knowledge that will become accessible with the 72

73 advent of trace element and isotope analytical tech-
 74 niques with higher spatial resolution. The contribution
 75 of Reid et al. provides good evidence for a direct link
 76 between rheological properties and diffusion at ele-
 77 vated pressures and temperatures. Reid et al. present
 78 in situ viscosity measurements of silicate liquid as a
 79 function of pressure, and show that liquid viscosities
 80 compare favourably with viscosity estimates derived
 81 from oxygen self-diffusion measurements in these
 82 liquids at the same conditions. Dobson presents elec-
 83 trical conductivity measurements in sodium-bearing
 84 Mg-silicate perovskite, the main phase of the Earth's
 85 lower mantle, at very high pressures and tempera-
 86 tures. He finds that the electrical conductivity in this
 87 phase is controlled by diffusion of oxygen vacancies
 88 at low temperatures—another illustration of the link
 89 between physical properties and diffusion—with a
 90 possible transition to intrinsic electronic conduction
 91 at high temperatures. Finally, Watson et al. provide
 92 diffusion rates for some highly siderophile elements
 93 in solid iron–nickel alloys. These data are directly rel-
 94 evant for estimates of the cooling rates of meteorites
 95 and their parent bodies, and could also be relevant to
 96 the evolution of planets with solid metallic cores.

97 Fortenfant et al. provide new experimental data on
 98 the partitioning of highly siderophile elements be-
 99 tween liquid metal and (Mg, Fe)O at pressures relevant
 100 to the differentiation of Earth's core and mantle. Data
 101 like these give tight constraints on early planetary ac-
 102 cretion models. The Allan et al. paper reviews recent
 103 progress in the relatively new field of computer sim-
 104 ulations of trace element partitioning. The tools de-
 105 scribed are likely to become more widely used in the
 106 near future, to study partitioning at extreme conditions
 107 not amenable to direct experimentation. As an illustra-
 108 tion of the power of these tools, Corgne et al. provide a
 109 comparison between experimental and computational
 110 data on trace element partitioning between Ca- and
 111 Mg-silicate perovskites and silicate melts. Because of
 112 the very high pressures (exceeding 23 GPa) and tem-
 113 peratures (exceeding 2500 K) required, systematic ex-
 114 periments involving these phases are still very chal-
 115 lenging, and insights from atomistic simulation will
 116 certainly aid in the interpretation of experimental re-
 117 sults. McDade et al. discuss new experimental data on
 118 mineral–melt trace element partitioning during melting
 119 of upper mantle material, relevant to the formation of
 120 mid-ocean ridge basalt. The paper is an example of re-

cent attempts to produce partition coefficients directly
 applicable to the modelling of complex petrological
 processes. Finally, Draper et al. present the most com-
 prehensive collection of data on majorite–melt trace el-
 element partitioning to date. By using a starting compo-
 sition similar to recent estimates of the Martian mantle
 they illustrate how mineral–melt trace element parti-
 tioning studies are becoming increasingly relevant to
 planetary science studies.

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